

Quantifying Carbon Cycle- Climate Feedbacks

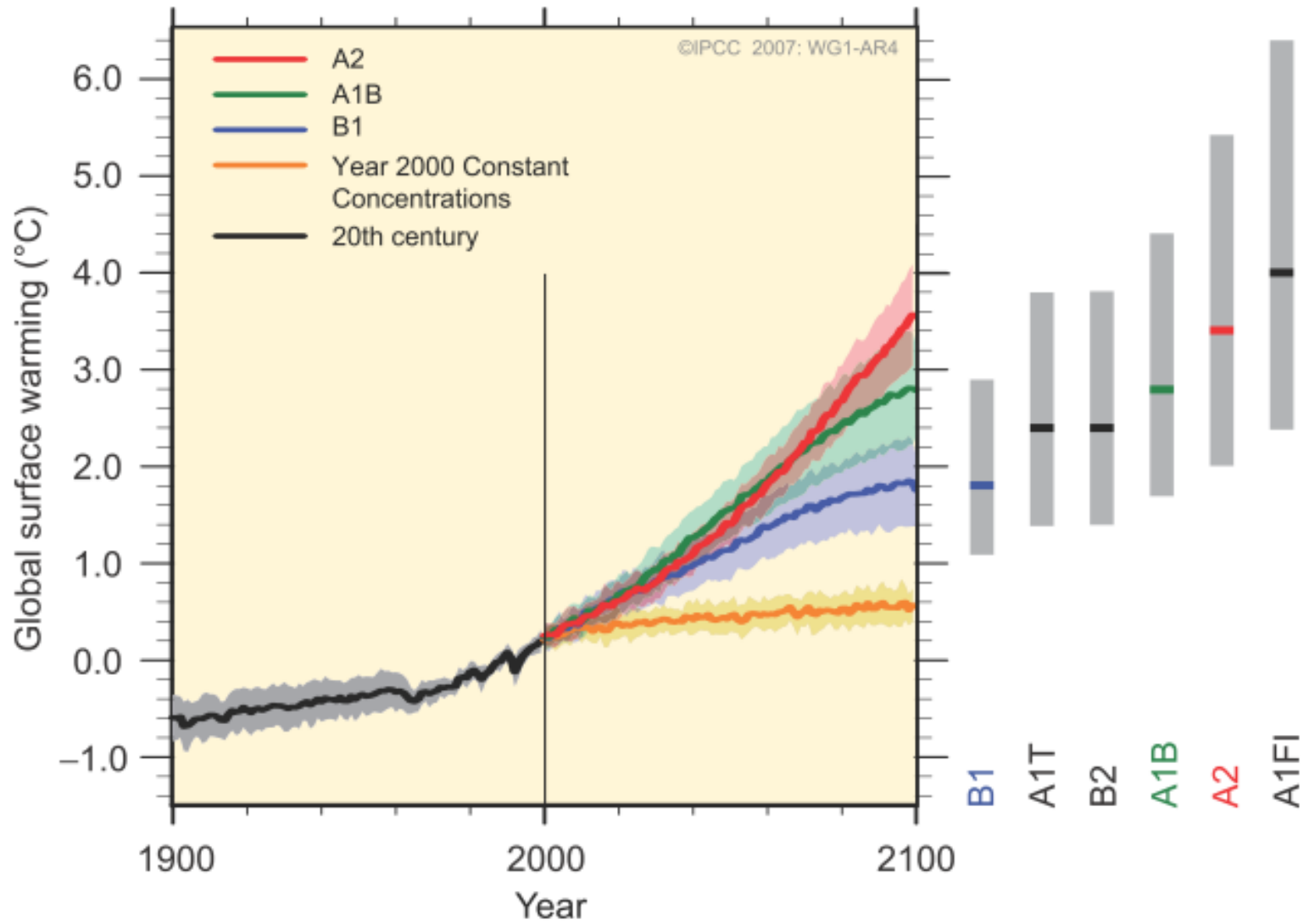
Jonathan Moch

Mentors: Thomas Froelicher, Keith
Rodgers, and Jorge Sarmiento

Feedbacks Overview

- $T_e = [S_o(1-A)/4\sigma]^{1/4} = (s/\sigma)^{1/4}$
- $\Delta T_{eq} = f \Delta t_o$
- Positive or negative
 - E.g.: Permafrost melt, CO₂ fertilization
- $\Delta T_{eq} = \Delta T_o + \Delta T_{feedbacks}$
- Objective: Deconstruct Carbon Cycle Portion of $\Delta T_{feedbacks}$
 - Useful for comparing models and getting at mechanisms

MULTI-MODEL AVERAGES AND ASSESSED RANGES FOR SURFACE WARMING



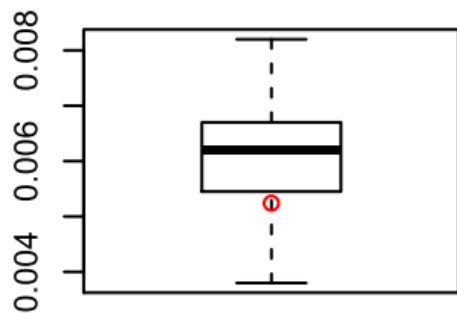
Terminology

- $\alpha = \text{K} / \text{ppm CO}_2$
 - Linear transient climate sensitivity
- $\beta = \text{GtC} / \text{ppm CO}_2$
 - Sensitivity of carbon uptake to atmospheric CO_2
- $\gamma = \text{GtC} / \text{K}$
 - Sensitivity of carbon uptake to temperature change
- $g = (\Delta T_{\text{feedbacks}}) / (\Delta T_{\text{eq}})$
 - Ratio of feedback response to total response

ESM2M Analysis

- Three runs with 1% increase CO₂ per year over preindustrial levels until point of doubling
 - Fully coupled, radiatively uncoupled, biogeochemical uncoupled
 - Performed linear regressions on different runs to calculate feedback factors
 - Grid cell by grid cell regressions to perform regional analysis

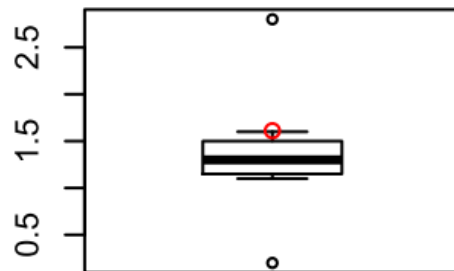
Alpha



0.00524

$$\Delta T^c = \alpha \Delta C_A^c$$

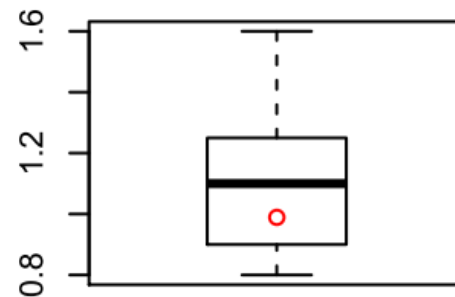
Beta Land



1.611129

$$\Delta C_L^u = \beta_L \Delta C_A^u$$

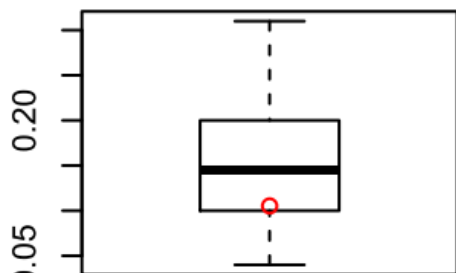
Beta Ocean



0.9890785

$$\Delta C_O^u = \beta_O \Delta C_A^u$$

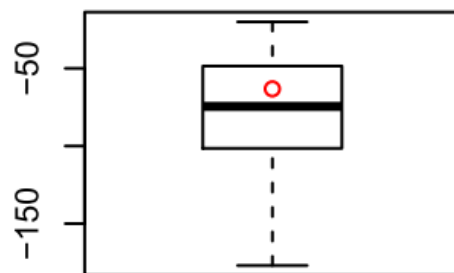
Gain



0.1052739

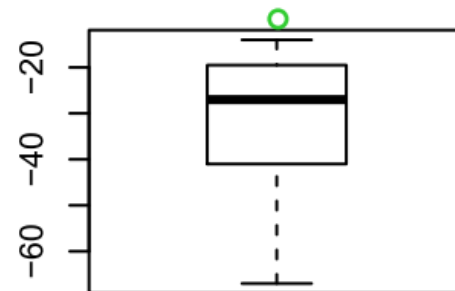
$$g = -\alpha(\gamma_L + \gamma_O)/(1 + \beta_L + \beta_O)$$

Gamma Land



-63.17894

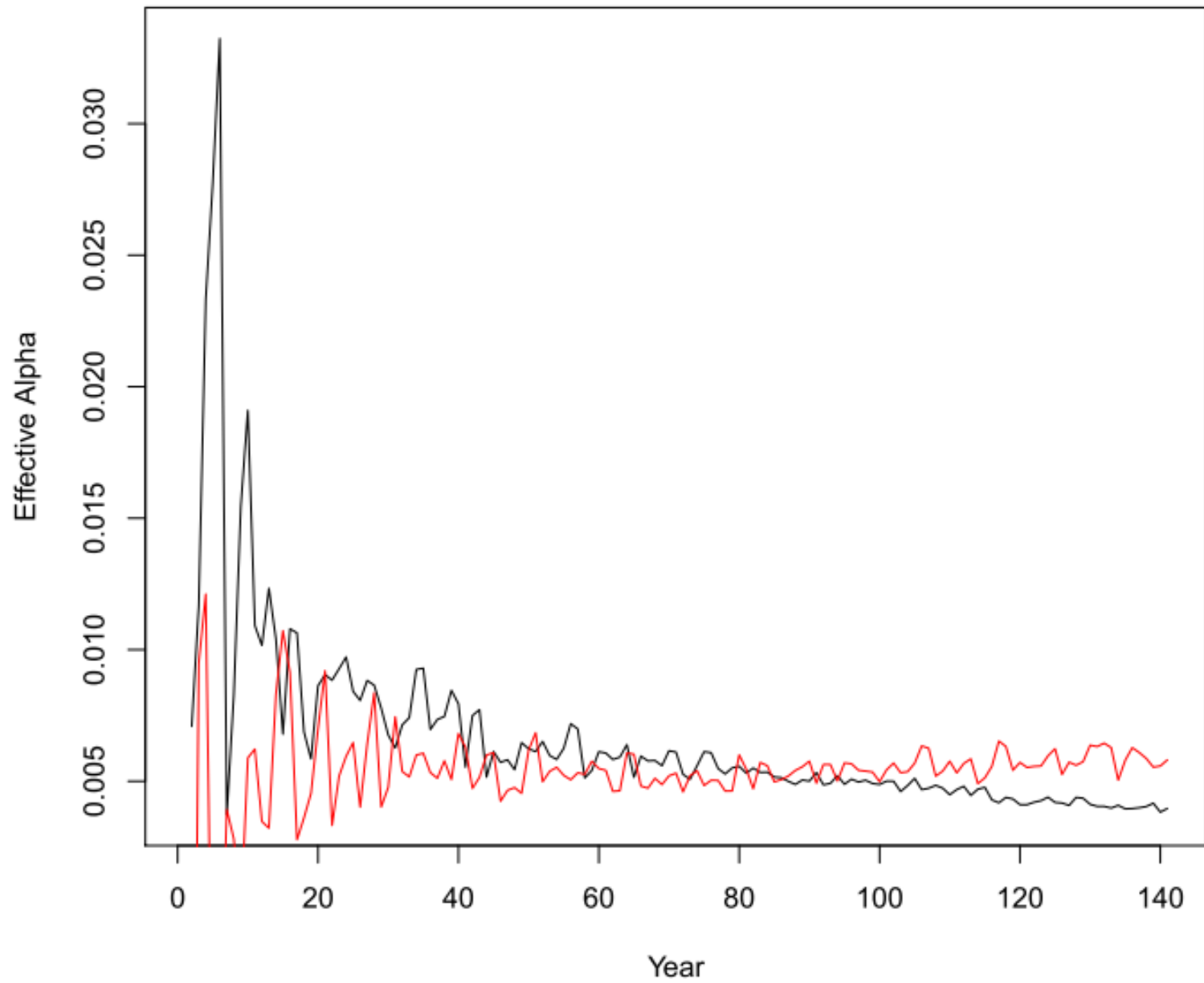
Gamma Ocean



-7.290276

$$\Delta C_O^c = \beta_O \Delta C_A^c + \gamma_O \Delta T^c$$

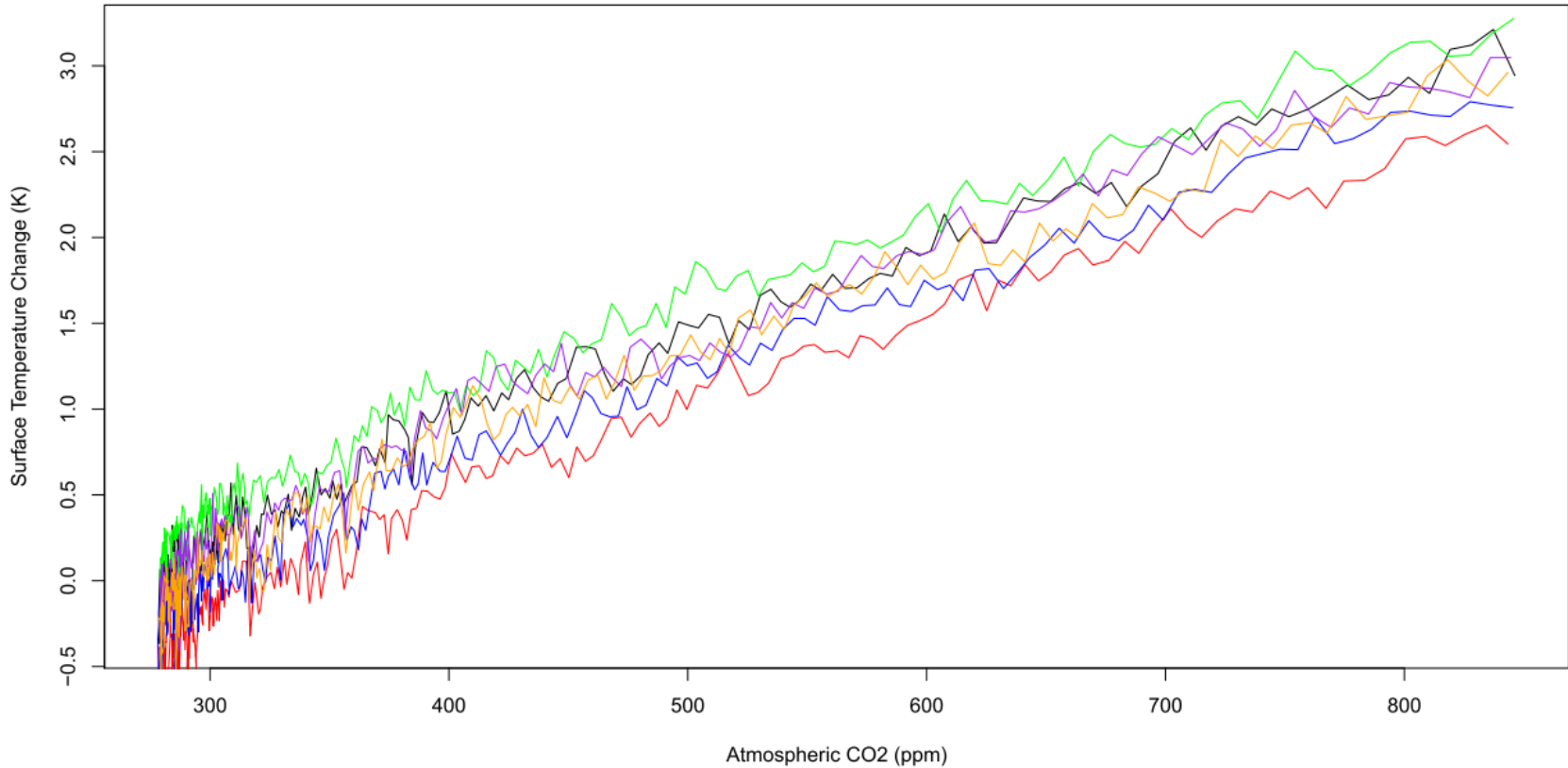
1% Concentration Scenario Alpha Progression



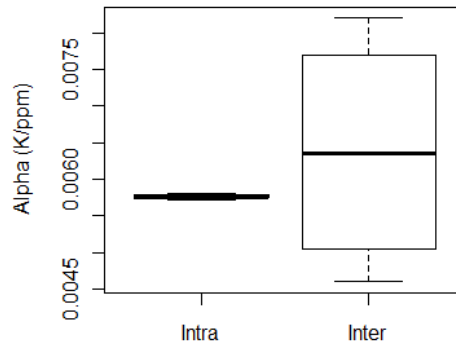
Ensemble Analysis

- NCAR CSM1.4
- 6 different ensemble fully coupled runs
- Used Friedlingstein et al 2006 method
- β calculated from one radiatively uncoupled run
- All runs from 1820 to 2000 using historical forcings, from 2000 to 2100 using A2 scenario

Simulated surface Temperature Response to Atmospheric CO2

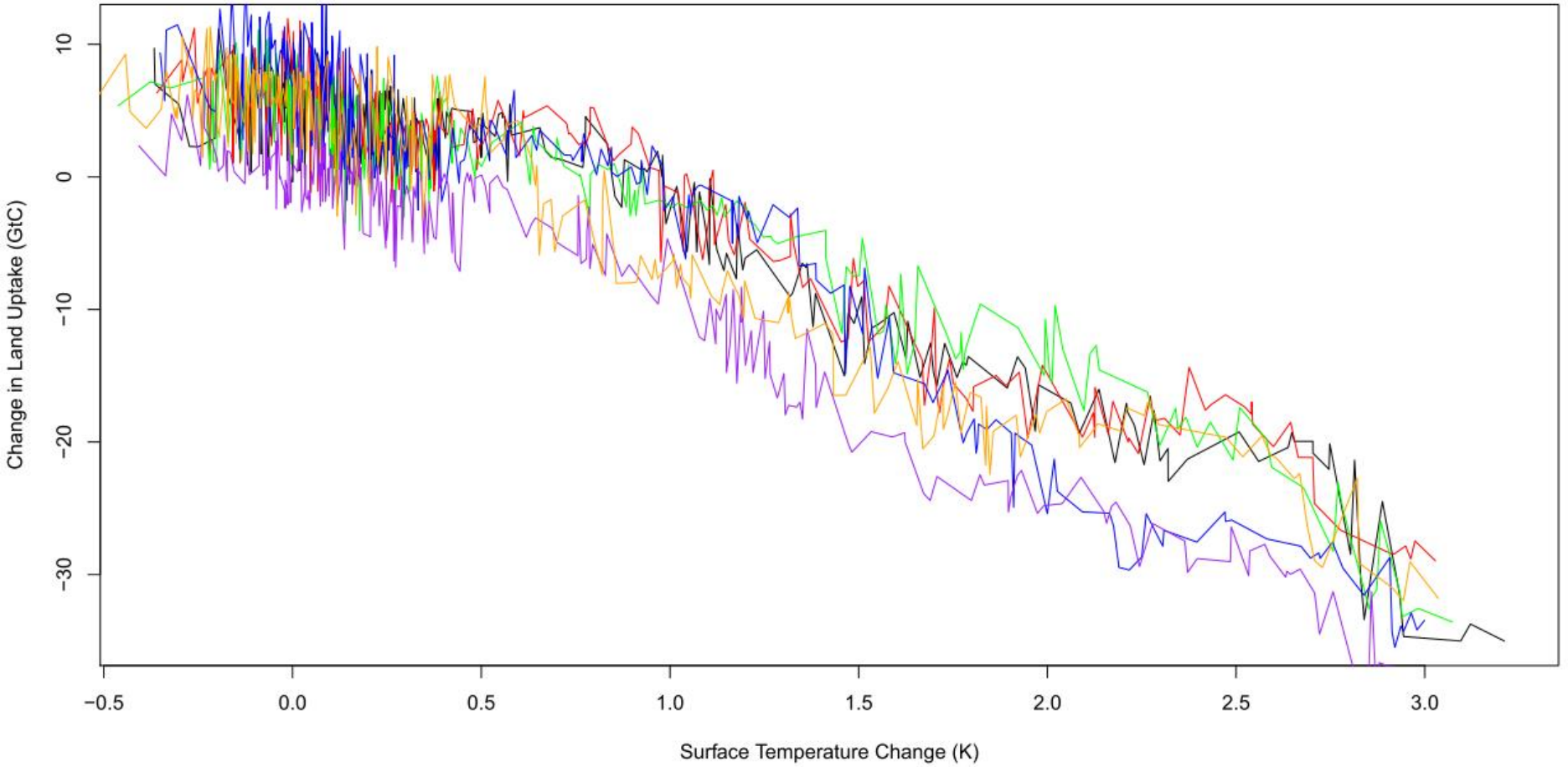


Intra-model	
NCAR Run	Alpha
a-1830	0.005786
b-1849	0.005746
c-1869	0.005746
d-1879	0.005721
e-1907	0.005770
f-pap	0.005811
Avg	0.005763
SD	0.000032321

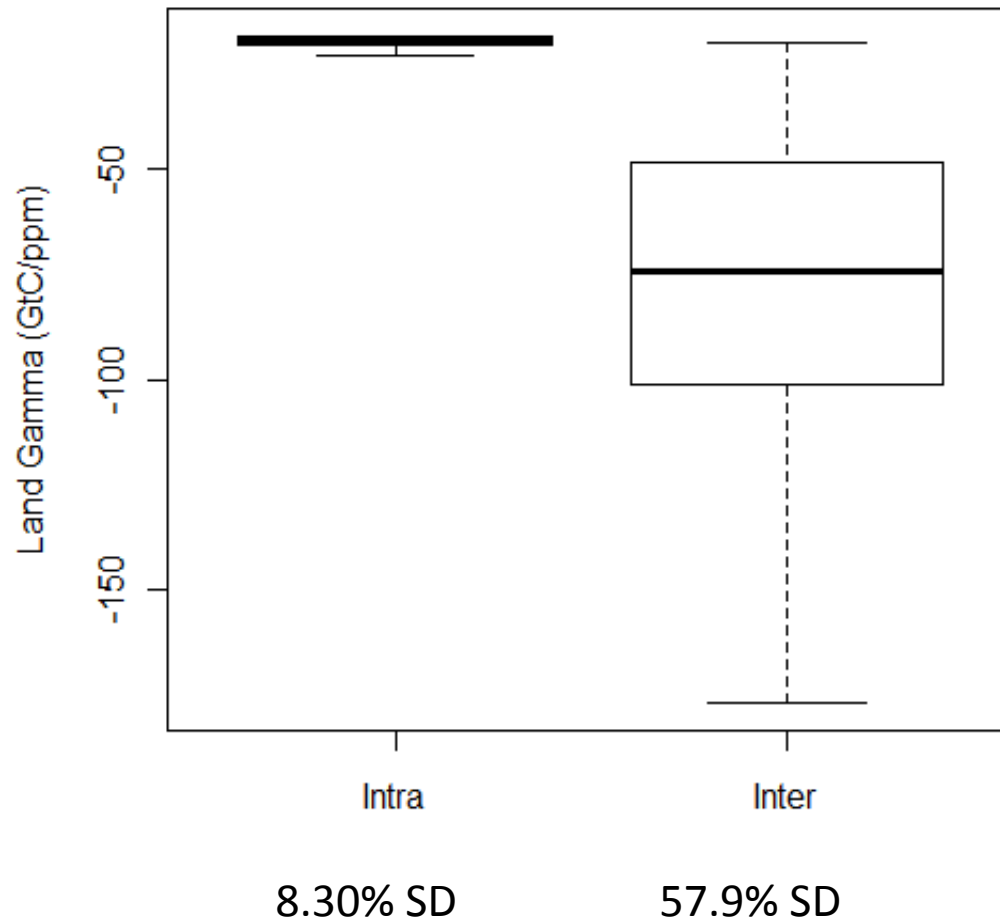


Inter-model	
Model	Alpha
IPSL	0.0072
NCAR	0.0046
BCM-C	0.0055
MPIM	0.0082
Avg	0.006375
SD	0.001626

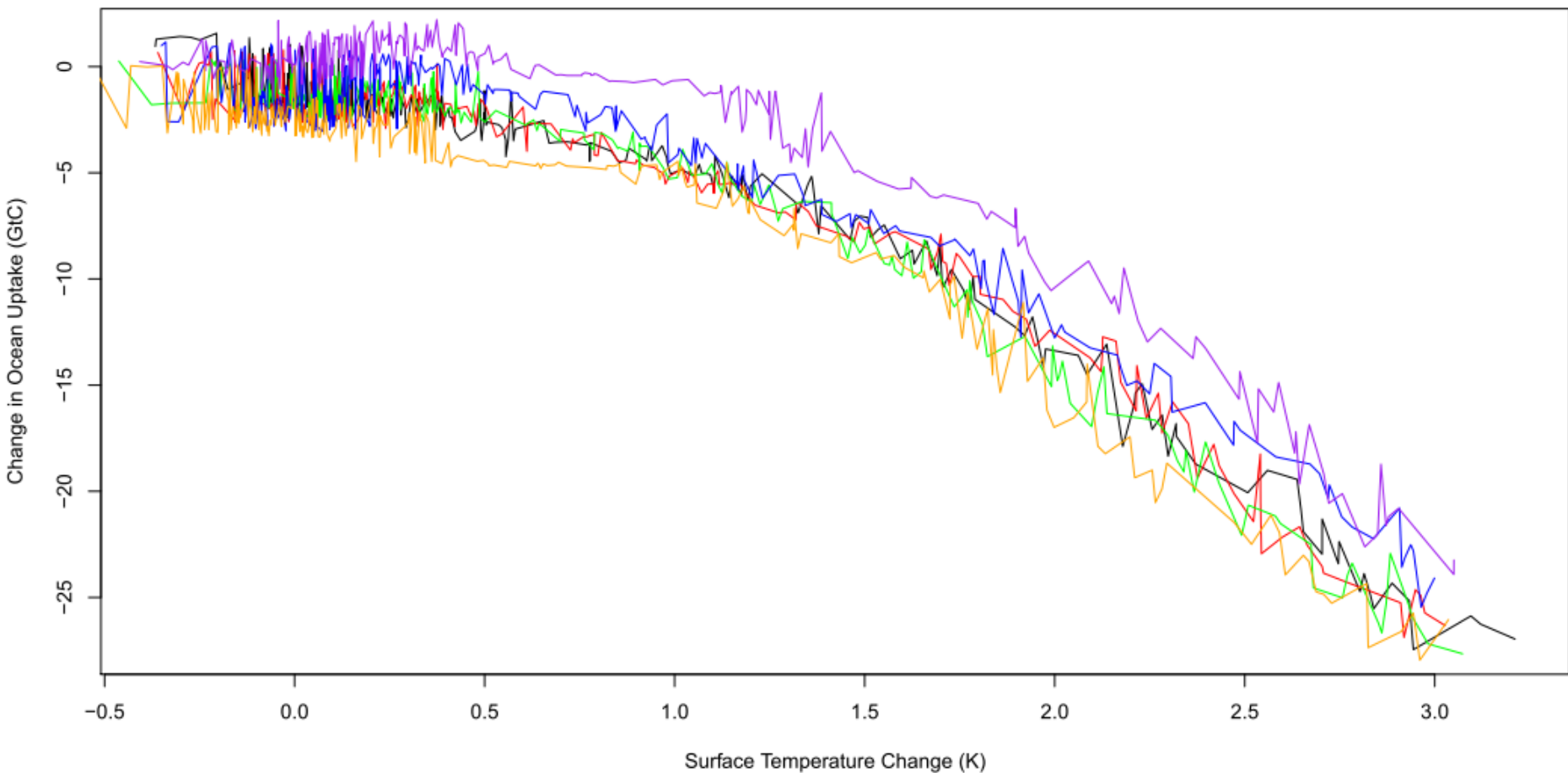
Sensitivity of Land Storage to Climate



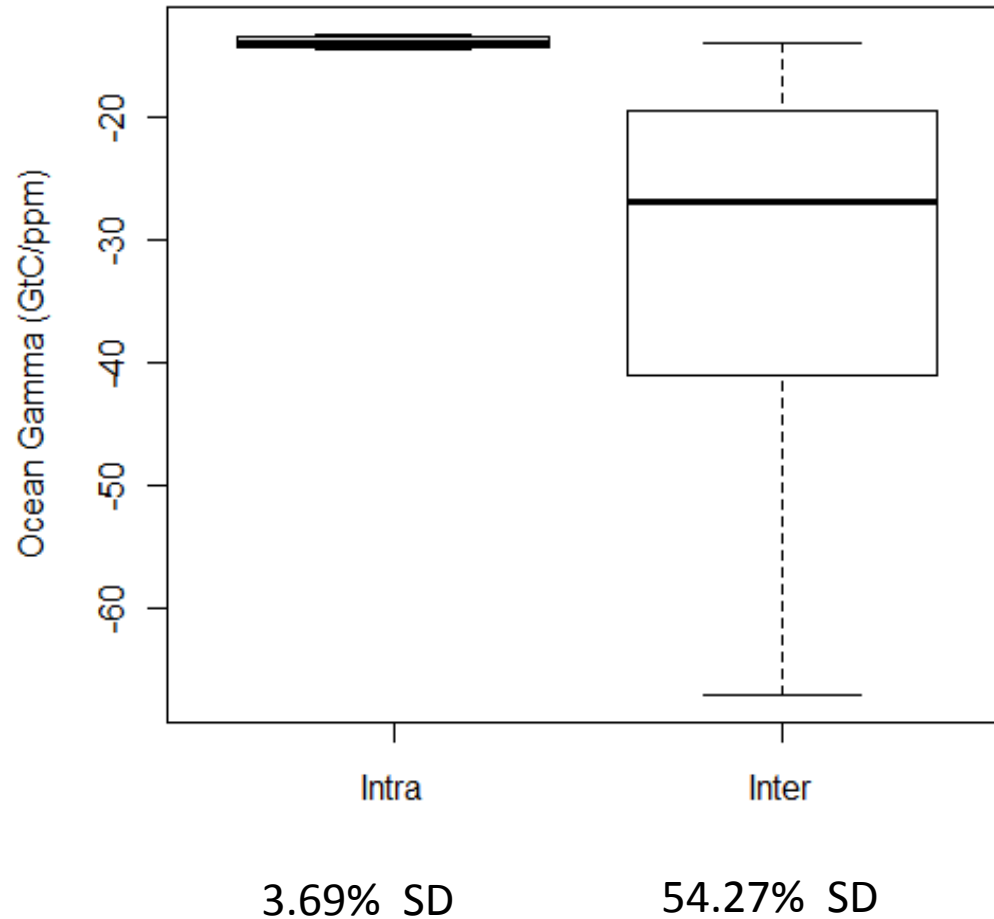
Intra & Inter Model Land Gamma



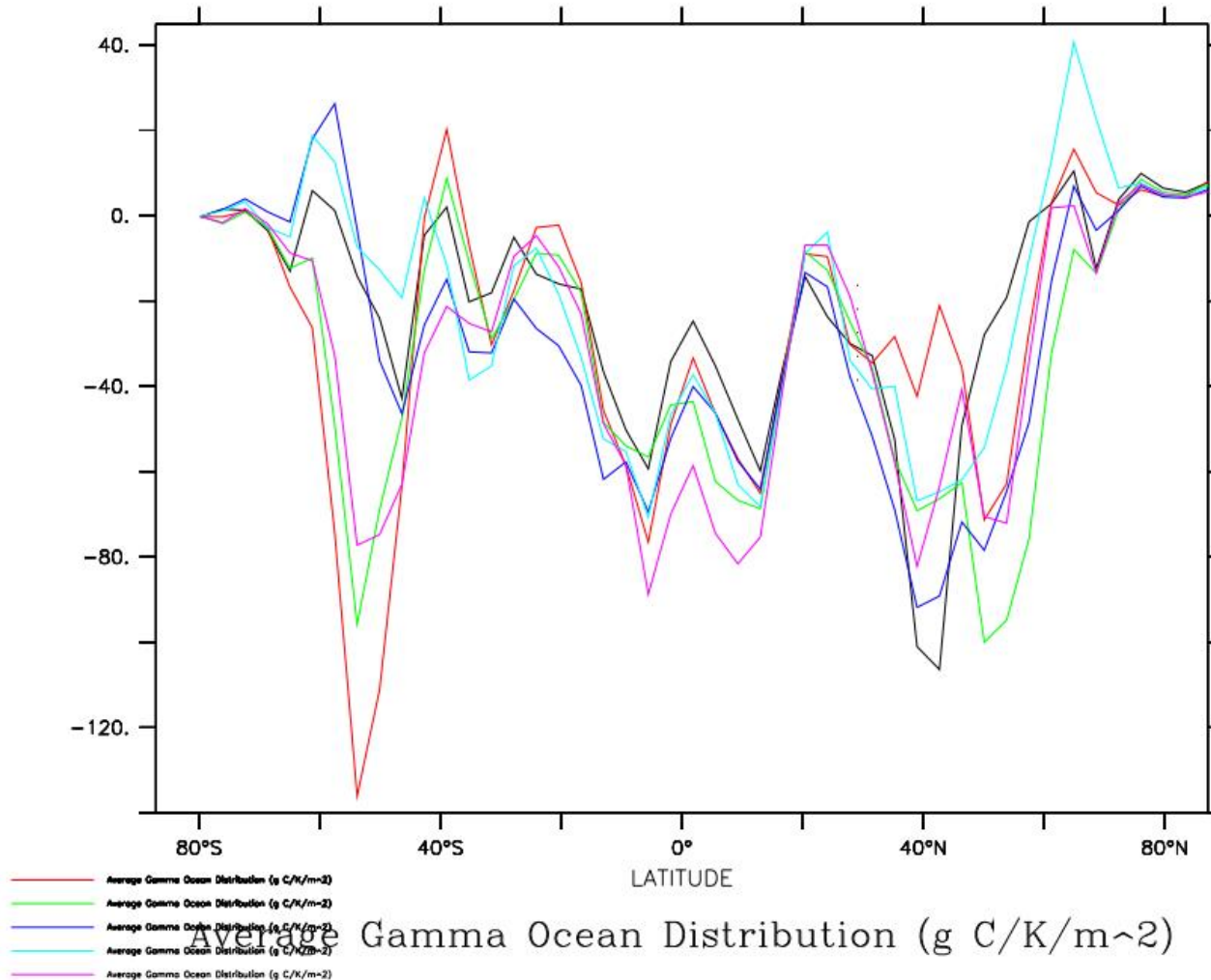
Sensitivity of Ocean Storage to Climate



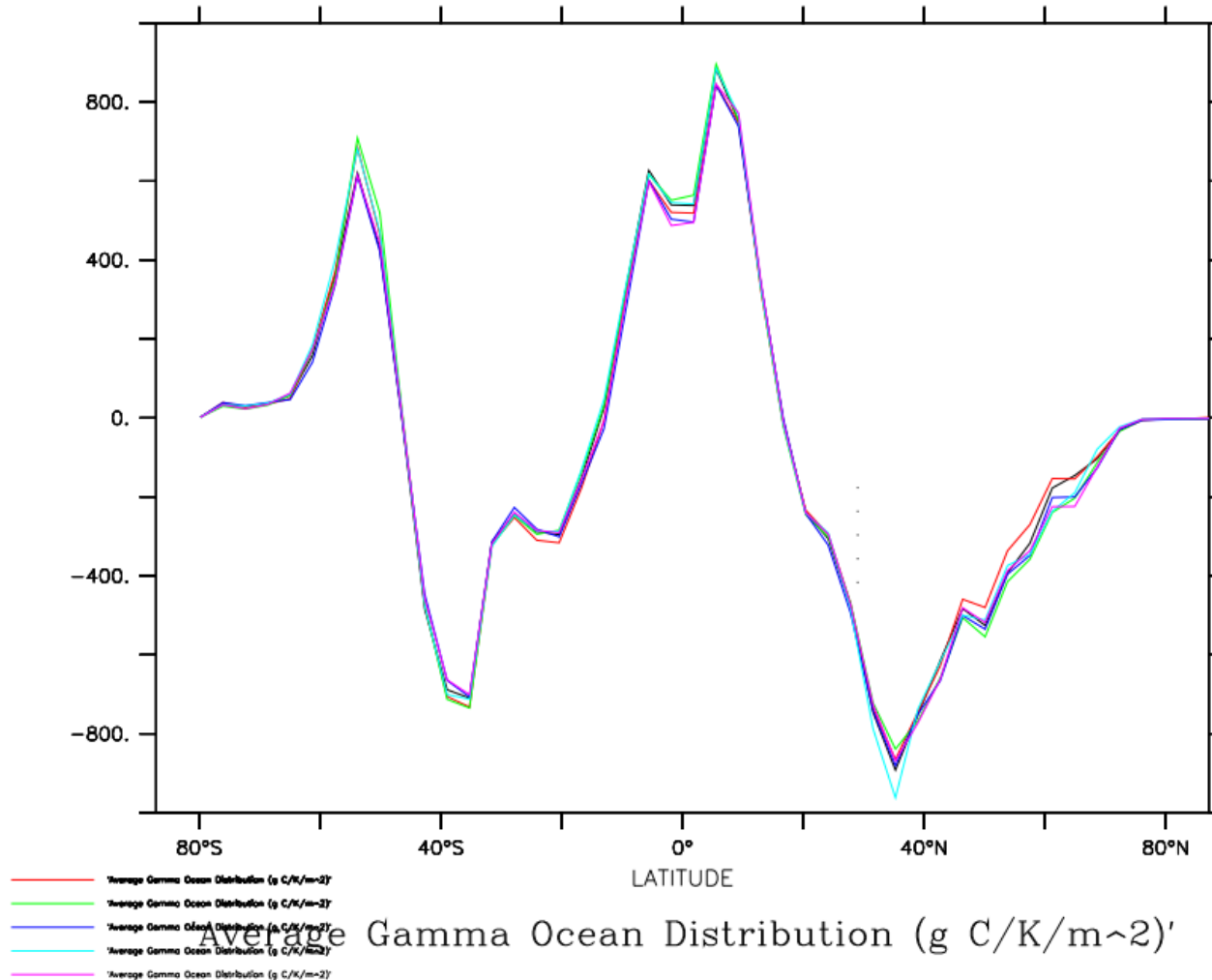
Intra & Inter Model Ocean Gamma



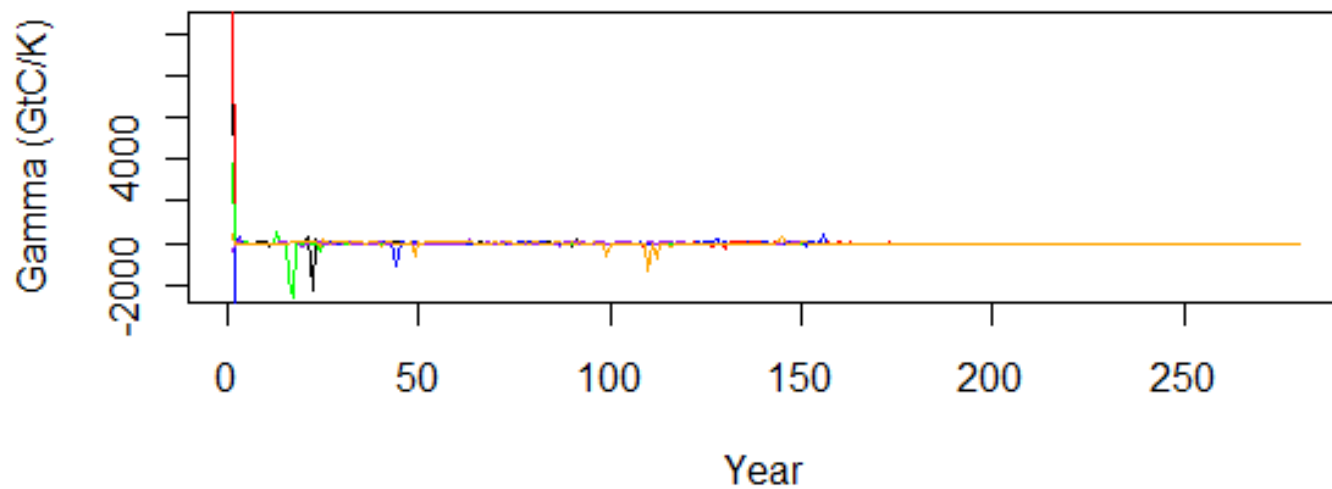
Ocean Gamma: 1820-2100



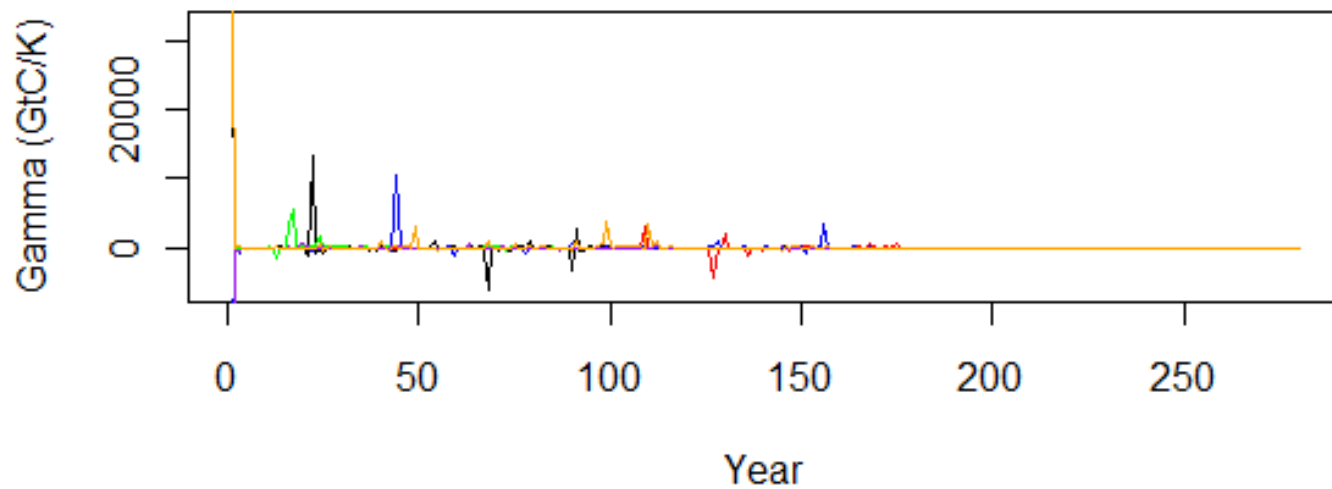
Ocean Gamma: 2010-2100



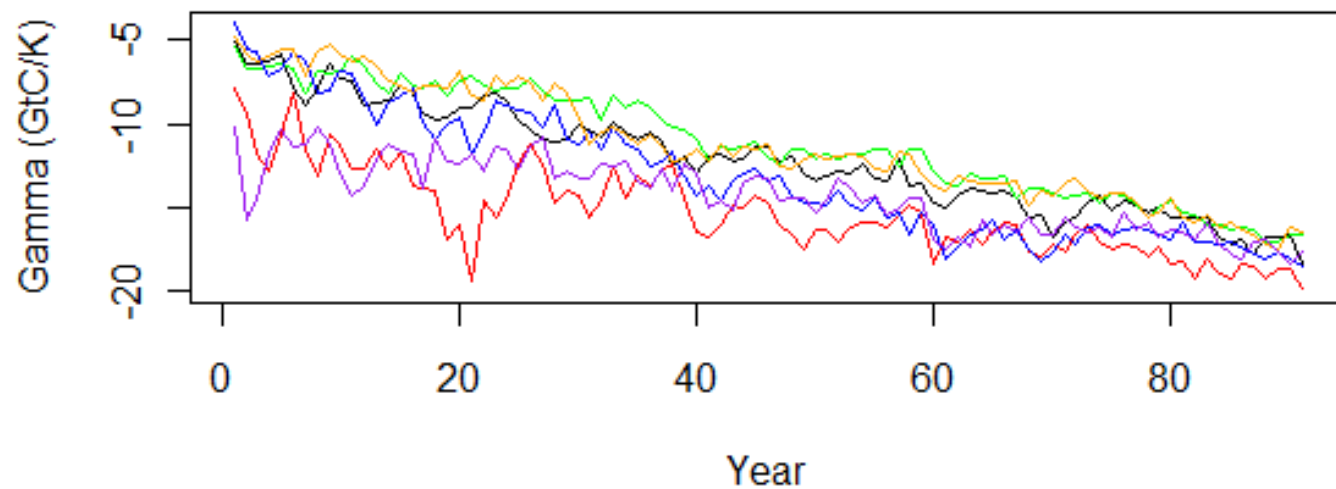
Ocean Gamma



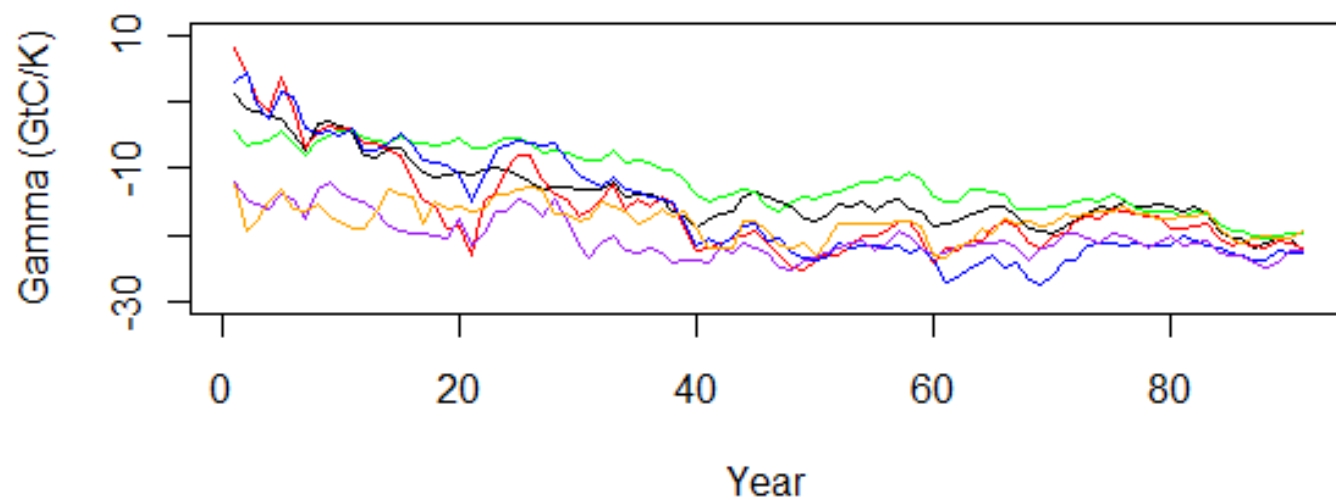
Land Gamma



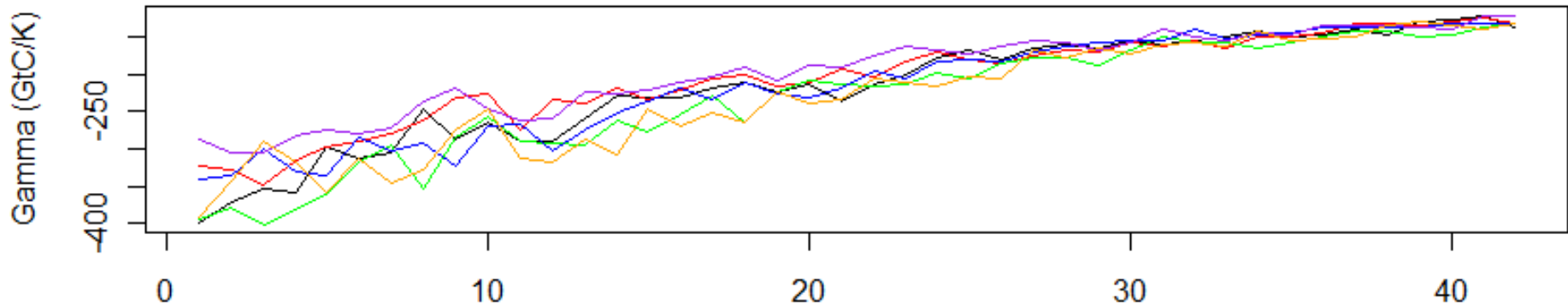
Ocean Gamma: 2010-2100



Land Gamma: 2010-2100

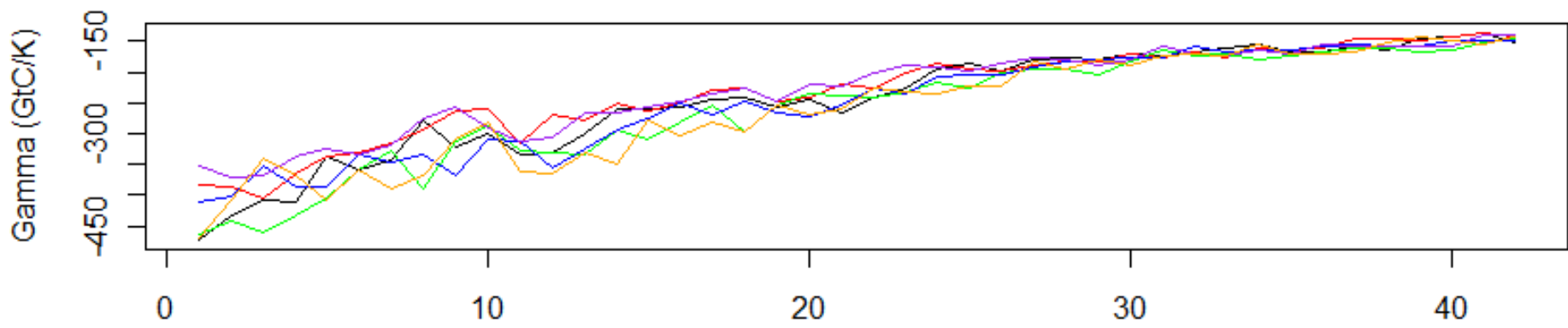


Ocean Gamma from 2010



Year	All	2010-2100
% SD	3.69	2.81

Land Gamma from 2010



Year	All	2010-2100
% SD	8.30	12.63

Conclusions

- ESM2M relatively unresponsive
- Feedback factors are not actually linear
- Regressions yield different results from instantaneous slope when started from different points
- Regional differences appear to converge after spin-up
 - Less clear on a global scale
- Larger β and γ if ignore spin-up, smaller α
- Intra-model uncertainty much smaller than inter-model uncertainty
 - Might make a difference on the margins

Thank You!